Sulfur Contents of Volcanic Rocks in Japan

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The total sulfur content was determined for 140 samples of Japanese volcanic rock by the mercury(II) thiocyanate photometric method of hydrogen sulfide formed by a wet reduction of the sample by tin(II)-strong phosphoric acid. Though determined values are widely scattered from 5390 down to 10 ppm (or less), the cumulative frequency distribution of the sulfur contents shows lognormality. The geometric and arithmetic means are 26 ppm (mg kg⁻¹) and 116 ppm, respectively. The alkaline rocks have high sulfur contents, the geometric mean of which is 62 ppm. Little significant difference was observed among the rhyolitic rock, the andesite and the trachyte, and the basaltic rock. Japanese and foreign standard rock samples were analyzed for sulfur. Almost all results were in good agreement with those obtained by other authors. In this case foreign standard samples also have higher sulfur contents than those of Japanese samples. The fact that the sulfur contents of Japanese volcanic rocks are lower than the previously accepted values seems to be characteristic of the islands arc of volcanism.

From the geochemical point of view sulfur is one of the most interesting and important volatile components of magma and is released as a volcanic gas into the atmosphere in the form of either sulfur dioxide or hydrogen sulfide. However, certain amounts are considered to be fixed in igneous rocks in such a form as sulfate or sulfide during the course of crystallization of magma.

Ever since the first work involving a determination of the sulfur in igneous rocks by Clark and Washington,¹⁾ many authors have reported sulfur contents in igneous rocks. Especially, from numerous samples Goldschmidt²⁾ has determined the average value of the sulfur contents in igneous rocks to be 520 ppm. However, his data, together with those of others, were widely scattered over more than one order of magnitude; thus, these determination methods seemed to lack reliability.

It was until the late 1950's when standard rock samples were prepared by geological surveys in several countries that the accuracies of individual determination methods could be compared with each other by using these standard samples. As a consequence, the reliability and accuracy of the determination method were much improved.3-14) Ricke7) has attempted a separate determination of the sulfate and sulfide sulfur in various kinds of igneous rocks; he found the total sulfur contents to be 300 ppm. From a series of such investigations he concluded that sulfur was generally inclined to be contained in alkaline rocks. He further pointed out that sulfide sulfur was more abundant in alkaline rocks whereas sulfate sulfur was rich in silicic rocks. Moore and Fabbi⁸⁾ reported the sulfur contents in deep ocean-floor erupted basalts to be 800±100 ppm. Recently Ueda and Sakai⁹⁾ pointed out that the sulfur contents of Quaternary volcanic rocks from the Japanese Islands Arc were less than 40 ppm, much lower than those of ocean-floor basalts.

The total sulfur contents of standard rock samples were summarized in review papers by Fleisher, 15)

Abbey,¹⁶⁾ and Flanagan.¹⁷⁾ All of these determined values were not always in good agreement among individual standard samples and fell, as a whole, in the range from several tens to several hundreds ppm. However, the total sulfur contents of the standard samples prepared by the Geological Survey of Japan showed with no doubt lower values by one order of magnitude than those of the foreign standard samples.

To find a reliable and feasible method of determining the total sulfur in rocks, we investigated and established several different methods by means of: (1) the photometric method, ¹⁸⁾ (2) the sulfide ion selective electrode method, ¹⁹⁾ and (3) the fluorometric method. ²⁰⁾ Each of these are based on a wet reduction of the sample by tin(II)-strong phosphoric acid. A fourth method, the UV-photometric method²¹⁾ is used after the combustion of a sample with vanadium pentaoxide. The determined values by these methods showed excellent agreement.

In this paper we present the total sulfur contents mainly determined by the photometric method¹⁸⁾ for 140 volcanic rocks in Japan.

Experimental

Detailed descriptions of each method for the determination of the total sulfur in volcanic rocks can be found in previous papers;^{18–21)} thus, only brief analytical description of the photometric method will be presented in this section.

Finely powdered rock samples are accurately weighed (100—300 mg) and heated with tin(II)-strong phosphoric acid at 280 °C for 15 min in a flow of carbon dioxide or nitrogen gas. Both the sulfate and sulfide sulfur contained in rock samples are separated from other components in a form of hydrogen sulfide. The hydrogen sulfide, thus formed, is absorbed in a zinc sulfate solution to be fixed as a zinc sulfide. In order to completely separate micro amounts of zinc sulfide by means of a centrifugal separator, a zinc hydroxide is coprecipitated by adding a sodium hydroxide solution to the absorbing solution. After separating the supernatant solution, an acetic acid, a mercury(II) thiocyanate-methyl alcohol solution and iron(III) nitrate-

perchloric acid solution are added to the precipitation. The mixture is shaken and diluted with redistilled water. Then, carbon tetrachloride is added and shaken vigorously to adsorb the suspending mercury(II) sulfide on the surface of small drops of carbon tetrachloride. After centrifugation, the absorbance of the supernatant solution is measured at the wavelength of 460 nm against the redistilled water as a reference. A standard sample to make a calibration curve is prepared as follows. Barium sulfate is mixed in an agate motor with a finely powdered obsidian in which no sulfur is assured to be detected by this method. The mixture is diluted with the obsidian several times until the sulfur content becomes 300 ppm. The lowest detectable limit of this method is 10 ppm for total sulfur. The coefficient of variation for five repeat analysis is 5% for a sulfur content of 200 ppm. The determination of the sulfur was performed in association with the main components and other elements such as halogen²²⁾ and boron²³⁾ for samples which were widely gathered from many locations in Japan by people listed in the acknowledgements.

Results and Discussion

Sulfur Contents of Volcanic Rocks in Japan. The total sulfur contents of volcanic rocks in Japan are summarized in Table 1. As should be apparent from Table 1, the determined values are widely scattered from the maximum value of 5390 ppm to the lowest detectable limit of 10 ppm. The sulfur contents of the volcanic rocks can be fairly well expressed as a lognormal type distribution when they are plotted on a logarithmic scale (log S ppm) on a cumulative frequency diagram (Fig. 1). Strictly, this curve seems to be composed of two straight lines with different slopes having a bend around log S=2.0. This feature can also be observed in other volatile components of volcanic rocks. For example, the cumulative

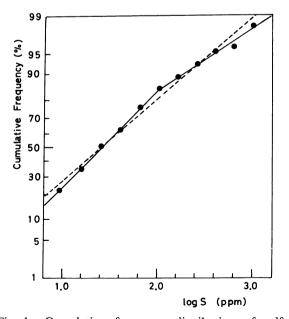


Fig. 1. Cumulative frequency distribution of sulfur contents of volcanic rocks.

frequency distribution of boron deviates from lognormality and the curve is composed of three straight lines; this is well understood by classifying rocks into three groups: alkaline rocks, glassy rocks and the other rocks.²³⁾ However, in the case of sulfur the bend in the curve could not be explained by such a classification into two groups, i.e., alkaline rocks and the rest other rocks, because both curves also showed a bend around $\log S=2.0$ (Fig. 1). The mean value of log S was estimated to be 1.41. The geometric mean of the total sulfur content is 26 ppm with a standard deviation of 0.64 ppm, while the arithmetic mean is This does not seem to reflect a good representative value of a statistical measure. It should be noted here that there are considerable numbers of samples that contain a total sulfur content less than this lowest limit. Threrefore, a more sensitive method is required in order to determine the exact distribution and geometric mean of the sulfur contents. Alkaline rocks have high sulfur contents, as pointed out by Ricke,7) and Ueda and Sakai.9) Their geometric mean is 62 ppm and can be clearly distinguished from those of other rocks by an inspection at the significant level 0.01. Moreover, the rocks from the Circum-Japan Sea Province have a geometric mean of the sulfur contents of 41 ppm, also showing a significant difference from the geometric mean of other rocks (at the significant In Fig. 2 are given the frequency distributions of each of the sulfur contents of three types of rocks classified by the silica content, i.e., rhyolitic rock, andesite and trachyte, and basaltic rock.

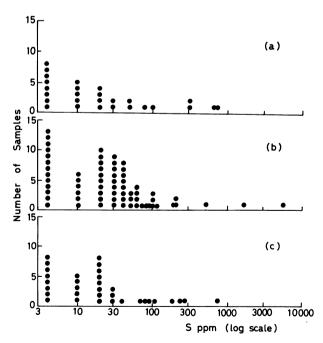


Fig. 2. Frequency distributions of sulfur contents of volcanic rocks.

(a): Rhyolitic rock, (b): andesite and trachyte, (c): basaltic rock.

Table 1. Sulfur Contents of Volcanic Rocks

No.	Sample	S/ppm	No	Sample	S/ppm
1	Basalt; Lava of 1950, Mihara-yama, O-shima,	90	38	Nepheline basalt; Nagahama, Hamada,	720
9	Tōkyō Metr.	20	20	Shimane Pref.	730
_	Basalt; Anei lava, Mihara-yama, Ö-shima	20	39	Dolerite; Kawai, Shidara, Aichi Pref.	190
3	Basalt; Okata, Ö-shima	40	40	Basaltic andesite; Dyke, Sukumo-gawa,	< 10
4	Basalt; Lava of 1962, Miyake-jima, Tōkyō	10	41	Hakone, Kanagawa Pref.	< 10
5	Metr.	10	42	Chilled margin of dyke No. 31 Andesite; Kami-futago-yama, Hakone	60
3	Basalt; Southern coast of Iga-tani, Miyake-	10	43	- '	
c	jima 660501-9	10	44	Andesite; Shimo-futago-yama, Hakone Andesite; Nagao-tōge, Hakone	<10
O	Basalt; Southern coast of Iga-tani, Miyake- jima 660501-10	10	45	Andesite; Nagao-toge, Hakone Andesite; Southeast of Okimura, Haha-jima,	<10
7	Basalt; Meiji-lava, Tosa Yuno-hama, Miyake-			Bonin Islands, Tōkyō Metr.	10
	jima	20	46	Andesite; Kitamuki-kakō lava, Tokachi-dake,	
8	Basalt; Kayakuko-ura, Miyake-jima	10		Hokkaido	100
9	Basalt; Oji-ike, Miyake jima	10	47	Andesite; Kami-horokametokku-kabu lava,	
10	Basalt; Wall of No. 3 entrance of the crater,			Tokachi-dake	40
	Miyake-jima	10	48	Andesite; Kumano-sawa-shita, Tokachi-dake	60
11	Basalt; Northern coast of Ako, Miyake-jima 660501-8-1	20	49	Andesite; Lava of parasitic crater (Minami- kakō), Yōtei-zan, Hokkaidō	<10
19	Basalt; Northern coast of Ako, Miyake-jima	20	50	Andesite; Somma lava, Usu, Hokkaidō	<10
14	660501-8-2	20	51	Andesite; Dome lava, Tarumai-yama,	_ 10
12	Basalt; Western foot of Higashi-yama,	40	31	Hokkaidō	30
13	Hachijō-jima Tōkyō Metr.	30	52	Andesite; Bomb, Tarumai-yama	30
14	Basalt; Southeastern foot of Nishi-yama,	30	53	Andesite; Eniwa-dake, Hokkaidō	10
17	Hachijō-jima	20	54	Andesite; Kenga-mine, Komaga-take,	10
15	Basalt; Sakasagawa, Itō. Shizuoka Pref.	10	34	Hokkaidō	20
	Basalt; Lava of 864, Aokiga-hara, Fuji-san,	10	55	Andesite; Around the crater, Komagatake	20
10	Yamanashi Pref.; No. 1	< 10		Andesite; Block in a mud flow,	20
17	Basalt; Lava of 864, Aokiga-hara, Fuji-san;	,		Komaga-take	20
	No. 2	< 10	57	Andesite; Lava of Yake-yama, Tamagawa	
18	Basalt; Kenmarubi lava, Fuji-yoshida, Fuji-			Spa, Akita Pref.	40
	san; No. 3	< 10	58	Andesite; Iwaya, Kakuda-yama, Niigata Pref.	20
19	Basalt; Maku-iwa, Fuji-san; No. 4	10	59	Andesite; Ejecta of Meiji eruption,	
	Basalt; Shirogane lava, Tokachi-dake,			Azuma-issaikyō-yama, Fukushima Pref.	30
	Hokkaidō	20	60	Andesite; Lava of 1783, Asama-yama,	
21	Basalt; Orimoto-tōge lava, Shidara, Aichi			Gunma Pref.	60
	Pref.	60	61	Andesite; Dyke, Kuniga, Nishino-shima,	
22	Basalt; Orimoto-tōge lava, Shidara	80		Oki, Shimane Pref.	80
23	Basalt; Orimoto-tōge lava, Shidara	210	62	Andesite; Iimi, Fuse, Dōgo, Oki	
24	Basalt; Kakuda-yama, Niigata Pref.	10		(epidote, pyrite bearing)	5390
25	Basalt; Nishimura, Nakamura, Dōgo, Oki,		63	Andesite; Chōshi, Harada, Dōgo, Oki	30
	Shimane Pref.	90	64	Andesite; Yunoo, Karatsu, Saga Pref.	< 10
26	Basalt; Kuniga, Nishino-shima, Dōdo, Oki	70	65	Andesite; Yakataga-udo, Neko-dake, Aso,	•
27	Basalt; Imazu, Saigō, Oki	30		Kumamoto Pref.	<10
28	Basalt; Kami-sue, Saigo, Dōgo, Oki	270	66	Andesite; South valley of Hinoo-toge,	•
29	Basalt; Chika-ishi, Harada, Dōgo, Oki	110		Neko-dake, Aso	20
30	Basaltic rock; No. 8 Shimo-motoya,	-	67	Andesite; Dyke, Ōshōji-iwa,	
	Nakamura, Dōgo, Oki	10		Yamaguchi-dani, Nekodake, Aso	30
31	Basalt; No. 15 Nishino-shima, Ura-gōri B,		68	Andesite; Chilled margin of dyke,	-
	Dōgo, Oki	80	-	Tengu-iwa, Neko-dake Aso	130
32	Basalt; Iwano, Karatsu, Saga Pref.; No. 1	20	69	Andesite; East valley of Jigoku-dani,	
33	Basalt; Iwano, Karatsu; No. 2	20		Neko-dake, Aso	10
34	Basalt; Taka-shima, Karatsu	20	70	Andeside; The 5th valley of western slope	-
35	Basalt; Hinode-matsu, Karatsu	<10		of Neko-dake, Aso	10
	Basalt; Hanase-zaki Kaimon-dake,	3	71	Andesite; Scoria lava of 1958,	
	Kagoshima Pref.	10		Neko-dake, Aso	50
37	Basalt; Trachyandesitic basalt; Intrusive		72	Andesite; Northern slope of Ōgi-yama, Aso	20
- •	sheet, South of Ōkuwa, Shidara, Aichi Pref.	240	73	Andesite; Sensui-yane, Taka-dake, Aso	30

Table 1. Continued.

		Table 1.	Contin		
No.	Sample	S/ppm	No.	Sample	S/ppm
74	Andesite; Northern slope of Narao-dake, Aso	20	107	Dacite; Usu-dake, Usu	20
75	Andesite; Southern slope of Naka-dake, Aso	30	108	Dacite; Dome lava, Ōusu-dake, Usu	20
76	Andesite; The top of Ogi-yama, Aso	10	109	Dacite; Komaga-take, Hokkaidō	20
77	Andesite; Southeastern slope of Naka-dake,		110	Dacite; Lava of somma, Akagi-yama,	
	Aso	10		Gunma Pref.	10
78	Andesite; Southwestern valley of		111	Dacite; Eboshi-dake, Aso, Kumamoto Pref.	320
	Naka-dake, Aso	20	112	Dacite; Koeboshi, Aso	20
79	Andesite; The top of Taka-dake, Aso	90	113	Rhyodacite; Eastern slope of Otogase	
80	Andesite; Eastern slope of Okamado-yama,			(Senriga-hama volcano)	30
	Aso	10	114	Liparite; Tenjō-san, Kōzu-shima, Tōkyō	
81	Andesite; Southern slope of Naka-dake, Aso	30		Metr.	10
82	Andesite; Road cutting near the Akamizu		115	Liparite; Southeastern foot of Tenjō-san,	
	station, Aso	20		Kōzu-shima	<10
83	Andesite; Kometsuka, Aso	< 10	116	Liparite; South of Naga-hama, Kōzu-shima	20
84	Andesite; Bomb, Eastern slope of Ōjō-dake,		117	Liparite; Takodo-yama, Kozu-shima	< 10
	Aso	30	118	Liparite; Mae-hama, Nii-jima, Tōkyō Metr.	10
85	Andesite; Crater wall of Ōjō-dake, Aso	100	119	Liparite; Nii-jima	10
86	Andesite; Northeastern Eboshi-dake,		120	Liparite; Southern coast of Shikine-jima,	
	Shigetama Spa, Aso	50		Tōkyō Metr.	10
87	Andesite; Northern slope of Naka-dake, Aso	10	121	Liparite; Southeastern coast of Shikine-jima	10
88	Andesite; Under the bridge on the		122	Liparite; Nakaya, Fuse, Dōgo, Oki,	
	Kuro-kawa, Aso	20		Shimane Pref.	660
89	Andesite; Lava of 1946, Sakura-jima,		123	Rhyolitic rock; Nishida, Saigō, Dōgo, Oki	50
	Kagoshima Pref.	190		Rhyolitic rock; Igo, Nakamura, Dōgo,	00
90	Andesite; Lava of 1914, Sakura-jima	200		Oki; No. 6	10
91	Andesite; Anei lava, Sakura-jima	110	125	Potash liparite; Manzō-yama, Shimoda,	
92	Andesite; Bunmei lava, Sakura-jima	40		Shizuoka Pref.	30
93	Andesite; Satsuma-iwō-jima, Kagoshima Pref.		126	Alkali rhyolite; Madara-jima, Saga Pref.	100
94	Andesite; Lava of Iwō-dake, inner cliff	, ,	127	Obsidian; Shirataki, Hokkaidō	<10
<i>3</i> 1	of Higashi Spa, Satsuma-iwō-jima	100	128	Obsidian; Wada-tōge (pale part)	<10
95	Andesite; Iwō-dake, Satsuma-iwō-jima	20	129	Obsidian; Koshi-dake, Imari, Saga Pref.	<10
96	Andesite; Inamura-dake, Satsuma-iwō-jima	40	130	Obsidian; Hime-shima, Ōita Pref.; No. 1	10
97	Andesite; Yahazu-dake lava, Sakamoto,		131	Obsidian; In a deposit of ejecta of Kikai	10
31	Satsuma-iwō-jima	50	131	Caldera, Satsuma-iwō-jima, Kagoshima Pref.	140
98	Andesite; The uppermost lava of Shin-dake,	30	132	Obsidian; Mifune	10
90	Kuchino-erabu-shima, Kagoshima Pref.	60	133	Trachyte; Harada, Saigō, Dōgo, Oki,	10
90	Glassy andesite; Southern slope of E-san,	00	133	Shimane Pref.	40
33	Hokkaidō	130	124	Trachyte; Funakoshi, Nishino-shima, Dōgo,	40
100		130	134		40
100	Glassy andesite; Shōwa-iwō-jima, Satsuma-iwō-jima, Kagoshima Pref.	70	125	Oki	40
101		70	135	Trachyte; Minato, Nakamura, Dōgo, Oki;	10
101	Glassy andesite; Shōwa-iwō-jima,	50	100	No. 5	10
100	Satsuma-iwō-jima (pumiceous part)	50	136	Trachyte; Uragōri, Nishino-shima, Dōgo,	1.000
102	Dacitic andesite; Sembe-yama, Shimane	- 10	107	Oki; No. 14—1	1600
100	Pref.	< 10	137	Trachyte; Uragōri, Nishino-shima, Dōgo,	200
103	Trachyandesite; Suribachi-yama, Iwō-jima,	E00	100	Oki; No. 14—2	200
104	Sulfur Islands, Tōkyō Metr.	500	138	Trachyte; Kojōga-hana, Saigō, Dōgo,	
104	Trachyte; Funakoshi, Nishino-shima, Oki,	40	100	Oki; No. 23	40
	Shimane Pref.	40	139	Peridotide; Horoman, Hokkaidō	10
105	Phonolitic trachyte; Tokage-iwa,	9.5	140	Syenite; Oyama, northeastern foot of	
	Tsuzurao-yama, Dōgo, Oki	30		Takuhi-yama, Nishino-shima, Oki,	
106	Dacite; Dome lava, Shōwa-shinzan, Usu,	222		Shimane Pref.	1550
	Hokkaidō	320			

Each geometric mean is 23 ppm (standard deviation σ =0.72 ppm) for rhyolitic rock, 29 ppm (σ =0.63 ppm) for andesite and trachyte, and 22 ppm (σ =0.61 ppm)

for basaltic rock, respectively. No significant difference could be observed among these three types of rocks. Though glassy rocks are enriched in such

Table 2. Comparison of Sulfur Contents of Rock Samples

Sample	Sulfur contents This work			ppm References	
•	a)	a) b) c)	d)		
W-1 Basalt	120				135^{7} , 124^{10} , 133^{11} , 113 ± 2^{12} 3^{26}
G-1 Granite	40				175^{7} , 74^{10} , 65^{11} , 3^{26}
JA-1 Andesite		21.4	22		16^{16} , 25 ± 5^{17}
JB-1 Basalt	20				13.1 ± 2.4^{12} , 14^{16} , 50^{27}
JB-2 Basalt		17.3	16		19^{16} , 23 ± 5^{17}
JB-3 Basalt		11.1			$10 \pm 3^{17)}$
JG-1 Granite	20				9.3 ± 0.2^{12}
JR-1 Rhyolite		11.7			15 ± 4^{17}
JR-2 Rhyolite		11.4			8±2 ¹⁷⁾
Andesite-1	200	198	200	200	
Andesite-2	110	103		110	

a) Photometric method. 18) b) Sulfide ion selective electrode method. 19) c) Fluorometric method. 20)

d) UV-photometric method.21)

volatile components as chlorine²²⁾ and boron,²³⁾ the sulfur content of every obsidian, a typical glassy rock, is lower than 10 ppm. This feature is consistent with the experimental result for the sulfur solubility in silicate melts by Katsura and Nagashima,24) i.e., sulfur has low solubility in silicate melts, especially in silica Moreover, sulfur seems to form the rich melts. heterogeneous phase in silicate rocks even at a fairly low concentration. The sulfur contents of volcanic rocks determined in this study are much lower than those of the igneous rocks reported before. Old data seem to lack reliability concerning the employed determination methods. However, in a recent review by Schneider,25) the values of the sulfur content of igneous rocks are not so different as the usual values, i.e., about 600 ppm. Samples in his paper are almost entirely from continental, ocean-floor or hot spots; very few are from islands arc. The low sulfur contents of Japanese volcanic rocks may be a characteristic of the islands arc volcanism, as pointed out by Ueda and Sakai.9) This feature is probably attributed (1) to the extent of an outgassing of sulfur during the course of magmatic emanation and solidification, (2) to the difference of the source magma of the islands arc from those of others, (3) to the isolation of the sulfur containing phase during the fractional crystallization of magma and (4) to the heterogeneous distribution of sulfur in the upper mantle. However, the fact that the sulfur contents of volcanic rocks of the islands arc are low should be discussed more carefully in correlation with the contents of other volatile components.

Comparison of Sulfur Contents of Standard Rock Samples. To confirm the reliability of our determination method, standard rock samples were analyzed by means of three different methods established by the present author; 18-20) the results were compared with those reported by other authors. The sulfur contents of two kinds of andesite were also

compared. The results are shown in Table 2. The present three methods (photometric method, 18) sulfide ion selective electrode method, 19) and fluorometric method²⁰⁾) are all based on a wet reduction by tin(II)-strong phosphoric acid. These methods together with the UV-photometric method gave good agreeable results for common andesite samples. The results for some standard samples and andesites by the three methods were also in good agreement with each other.

The sulfur content of standard rock sample G-1 (determined by the present author) showed fairly good agreement with those obtained by Frost and Thomas, 10) and Murphy and Sergeant 11) (ranging from 40 to 74 ppm). For the sulfur in W-1, a satisfactory agreement was also obtained between our result and those of Ricke,7) Frost and Thomas,10) Murphy and Sergeant,11) and Ueda and Sakai12) (showing contents between 120 and 135 ppm). The results of the present author for the standard rock samples of Japan were in good agreement with those by Ueda and Sakai, 12) Terashima, 13) and Terashima et al. 14) (about 20 ppm); this is very close to the geometric mean (26 ppm) of the volcanic rocks in Japan. The sulfur contents of Japanese standard rock samples also show lower values than those of foreign standard rock samples.

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References

- 1) R. W. Clark and H. S. Washington, U. S. Geol. Surv. Prof. Paper, 1924, 127.
- 2) V. M. Goldschmidt, Skr. Noeske Videns.-Akad. Oslo, I. Mat.-naturv. Kl., No. 4 (1937).
- 3) M. E. Coller and R. K. Leininger, *Anal. Chem.*, **27**, 949 (1955).
- 4) R. P. Larsen, L. E. Ross, and N. B. Ingber, *Anal. Chem.*, **31**, 1956 (1959).
 - 5) J. G. Sen Gupta, Anal. Chem., 35, 1971 (1963).
- 6) A. D. Wilson, G. A. Sergeant, and L. J. Linnel, *Analyst*, **88**, 138 (1963).
 - 7) W. Ricke, Geochim. Cosmochim. Acta, 21, 35 (1960).
- 8) J. G. Moore and B. P. Fabbi, Contrib., Mineral. Petrol., **33**, 118 (1971).
- 9) A. Ueda and H. Sakai, *Geochim., Cosmochim. Acta*, **48**, 1837 (1984).
- 10) I. C. Frost and J. A. Thomas, cited in M. Fleischer, Geochim. Cosmochim. Acta, 29, 1263 (1965).
- 11) J. M. Murphy and G. A. Sergeant, *Analyst*, **99**, 515 (1974).
- 12) A. Ueda and H. Sakai, Geochemical J. 17, 185 (1983).
- 13) S. Terashima, Bull. Geol. Surv. Jpn, 30, 609 (1979).

- 14) S. Terashima, T. Yamashige, and A. Ando, *Bull. Geol. Surv. Jpn.*, **35**, 171 (1984).
- 15) M. Fleischer, Geochim. Cosmochim. Acta, 33, 65, (1969).
- 16) S. Abbey, Can. Spectrosc., 15, 10 (1970).
- 17) F. J. Flanagan, Geochim. Cosmochim. Acta, 37, 1189 (1973).
- 18) Y. Arikawa, T. Ozawa, and I. Iwasaki, *Bunseki*, *Kagaku*, **21**, 920 (1972).
- 19) Y. Arikawa, Nippon Kagaku Kaishi, 1986, 1074.
- 20) Y. Arikawa and H. Kawai, *Bunseki Kagaku*, **35**, 720 (1986).
- 21) Y. Arikawa, T. Ozawa, and I. Iwasaki, Bunseki Kagaku, 24, 497 (1975).
- 22) M. Yoshida, K. Takahashi, N. Yonehara, Y. Ozawa, and I. Iwasaki, Bull. Chem. Soc. Jpn., 44, 1844 (1971).
- 23) A. Isozaki, M. Yoshida, S. Utsumi, and I. Iwasaki, Nippon Kagaku Kaishi, 1973, 1896.
- 24) T. Katsura and S. Nagashima, Geochim. Cosmochim. Acta, 38, 517 (1974).
- 25) A. Schneider, "Hand Book of Geochemistry," Springer-Verlag, Berlin (1978), Vol. 2, Part 2, 16-E.
- 26) R. Brown and W. A. Wolstenholme, *Nature*, **201**, 598 (1964).
- 27) J. L. Bouvier, Private Communication, Geol. Surv. Canada (1970).